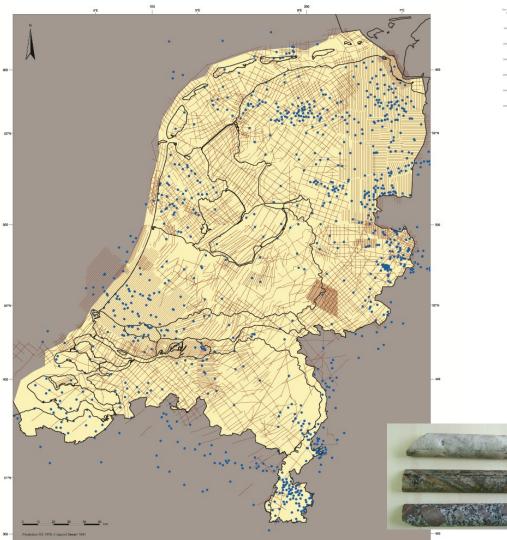


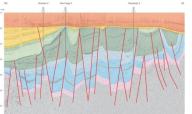
Geothermal exploration and geophysical techniques for sediment settings

- > Example by dutch aquifers
- Mature basin setting
- > How to use oil and gas data to find good aquifers
 - Seismic exploration
 - > Well logging techniques

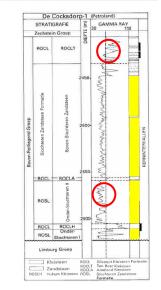


Dutch database: over 50 billion Euro of data





Well & Seismic Data Wells: 5876 Seismic: 72.000 km



<u>Log data</u> Gamma ray Sonic Resistivity Neutron, etc

Petrophysics Cores: 100 km Poro/perm: 60.000 measurements (300.000 total)



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ENERGY

OG NL Olie- en Gasportaal	Seismic surveys La Fields Aa Production Fa Infrastructure Sa	ublications and Data sets Links egislation Home dministrative procedures Disclaimer ees, taxes and state participation Contact eismicity and subsidence In het Nederlands) ontacts In het Nederlands)			
	Welcome to the NL Oil and Gas Portal				
This site provides information about oil and gas exploration and production in the Netherlands and the Dutch sector of		Recent changes			
	the North Sea continental shelf.	We keep this site continually up-to-date. Click <u>here</u> for an overview of recent changes.			
In land	It aims to help users access information furnished by the Dutch government in an easy, comprehensible fashion.				
		Other topics			

This site was produced at the request of the Dutch Ministry of Economic Affairs, Agriculture and Innovation and is being managed by TNO, Geological Survey of the Netherlands.

Other topics

Salt production

Underground gas storage

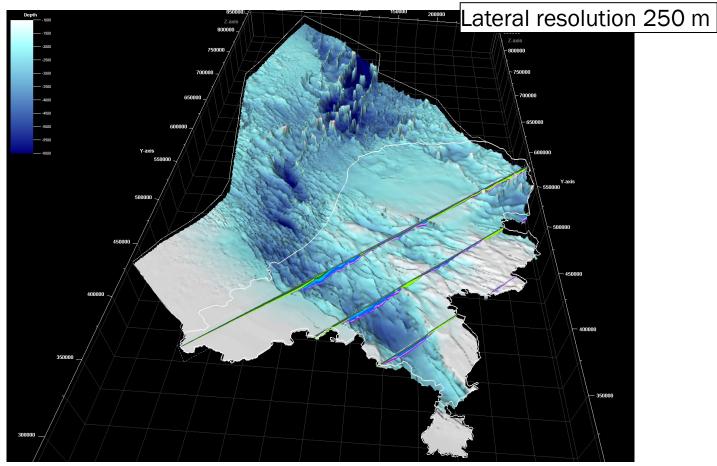
Geothermal Energy

Geological storage of CO2

•All kinds of well data & seismic data accessible and free to download



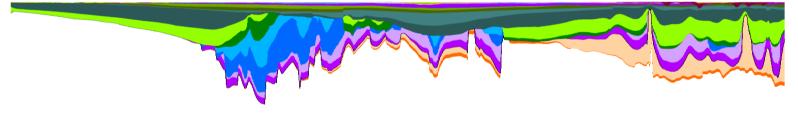
Seismic interpretation



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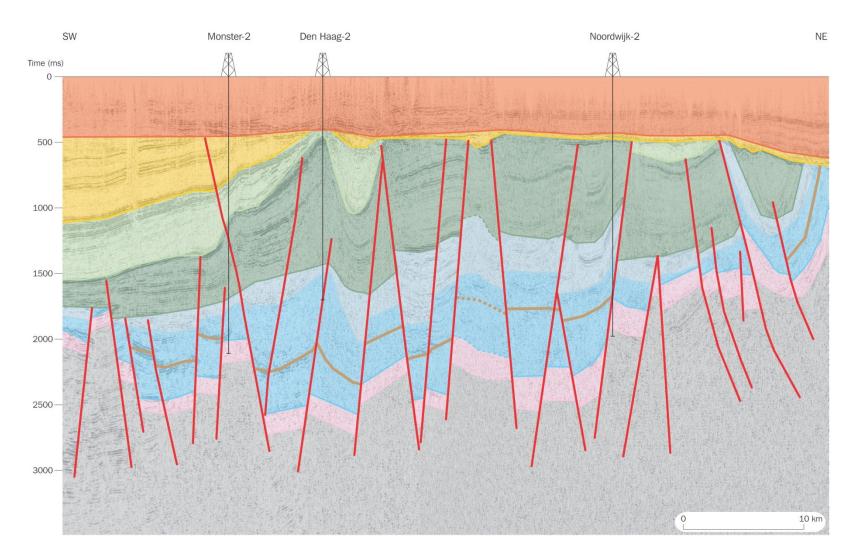
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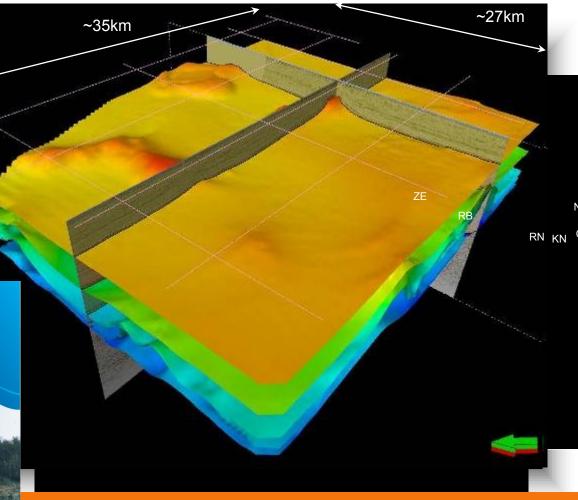
TNO innovation for life

Seismic reflection subsurface imaging



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3D Interpretation of seismic horizons

Nathantianda Institute of Applical Geoscience TNO - National Geological Survey

Geological Atlas of the Subsurface of the Netherlands - onshore



TNO 3D mapping of the subsurface

1985-2004





These methods can be divided into two main subclasses:

active seismic methods, which cover all seismic prospecting having an artificial sonic wave source;

passive seismic methods, which deal with the effects of natural earthquakes or those induced by fracturing related to geothermal fluid extraction and injection.

Seismic methods determine subsurface elastic properties influencing the propagation velocity of elastic waves and can be very helpful in obtaining structural information of the subsurface or even to outline a potential reservoir.



Elastic waves

When a stress is applied (or released) the corresponding strain propagates out from the source.

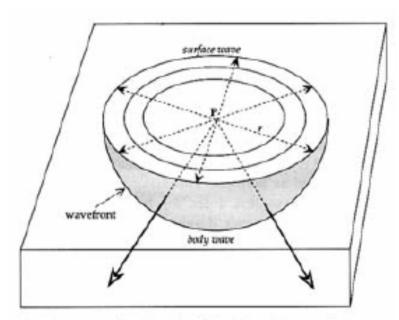


Fig. 3.9 Propagation of a seismic disturbance from a point source P near the surface of a homogeneous medium; the disturbance travels as a body wave through the medium and as a surface wave along the free surface. Point source seismic disturbance:

- Wavefront expands out from the point: Huygen's Principle
- Body waves: sphere
- Surface waves: circle
- Rays: perpendicular to wavefront

P-Waves

SEISMIC METHODS

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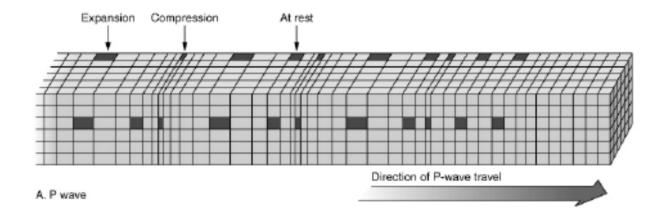
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- P for "primary" or "push-pull"
- Compression and rarefaction, no rotation
- Causes volume change as the wave propagates

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Similar to sound waves traveling through air



Can travel through solids/fluids/air



Body waves

S-waves

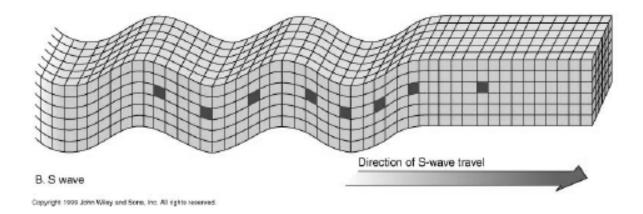
SEISMIC METHODS

ENERGY

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- S for "secondary" or "shear" and "shake"
- Shearing and rotation
- No volume change as the wave propagates

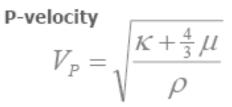


Cannot travel through fluids andair

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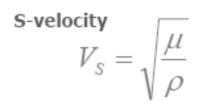
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P and S-velocities



change of shape and volume





change of shape only

For liquids and gases $\mu = 0$, therefore

 \rightarrow V_S = 0 and V_p is reduced in liquids and gases

➔ Highly fractured or porous rocks have significantly reduced V_p

The bulk modulus, κ is always positive, therefore $V_S < V_P$ always

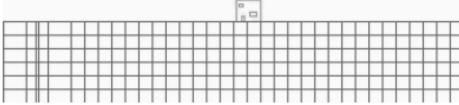
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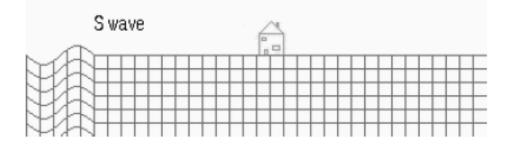
P-waves are the most important for controlled source seismology

- They arrive first making them easier to observe
- It is difficult to create a shear source, explosions are compressional



Body waves P and S-waves Arms P wave



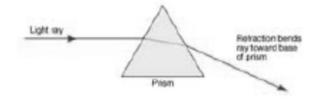


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Reflection and transmission

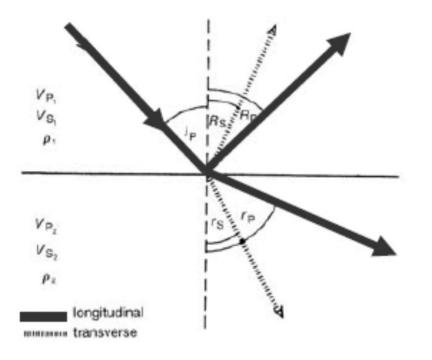
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Seismic rays obey Snell's Law

I E R G Y

(just like in optics)

The angle of incidence equals the angle of reflection, and the angle of transmission is related to the angle of incidence through the velocity ratio.

$$\frac{\sin i_P}{V_{P1}} = \frac{\sin R_P}{V_{P1}} = \frac{\sin r_P}{V_{P2}}$$

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Amplitudes reflected and transmitted

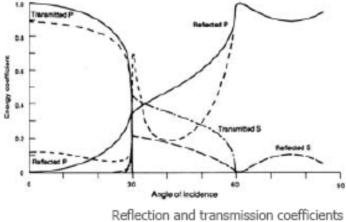
The amplitude of the reflected, transmitted and converted phases can be calculated as a function of the incidence angle using Zoeppritz's equations.

Simple case: Normal incidence Reflection coefficient

$$R_{C} = \frac{A_{R}}{A_{i}} = \frac{\rho_{2}V_{2} - \rho_{1}V_{1}}{\rho_{2}V_{2} + \rho_{1}V_{1}}$$

Transmission coefficient

$$T_{C} = \frac{A_{T}}{A_{i}} = 1 - R_{C} = \frac{2\rho_{1}V_{1}}{\rho_{2}V_{2} + \rho_{1}V_{1}}$$



for a specific impedance contrast

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These coefficients are determined by from the product of velocity and density – the impedance of the material.

 R_{C} usually small – typically 1% of energy is reflected.

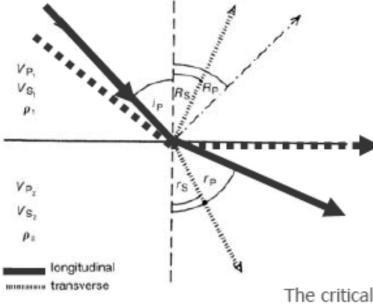
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SEISMIC METHODS

Critical incidence



 $\frac{\sin i_p}{V_{p_1}} = \frac{\sin r_p}{V_{p_2}}$

when $V_2 > V_1$, $r_p > i_p$

therefore, we can increase i_p until $r_p = 90^{\circ}$

When $r_p = 90^\circ i_p = i_C$ the critical angle

$$\sin i_C = \frac{V_{P1}}{V_{P2}}$$

The critically refracted energy travels along the velocity interface at V_2 continually refracting energy back into the upper medium at an angle i_c

→ a head wave

BASIS for refraction seismic used to detect boundaries of velocity contrasts





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ACTIVE SEISMIC METHODS

Seismic refraction surveys have been used to a limited extent because of the amount of effort required to obtain refraction profiles giving information at depths of 5 to 10 km, and the problems caused by the generally high degree of complexity of geological structures in areas likely to host geothermal systems.

Seismic refraction is normally restricted to cases where the densities of the rocks and thus seismic velocities increase with depth. In addition, geophone arrays for refraction measurements need a length of at least 4to 5 times (sometimes even 8 times) the sampling depth because of the very nature of refraction. The length requires higher shot energy (i.e., more explosives) and limits the applicability of refraction methods in exploration to shallower targets or to large-scale investigations of Earth's crust and upper mantle. Sometimes it can be used to get a first approximation about the velocity distribution at depth.

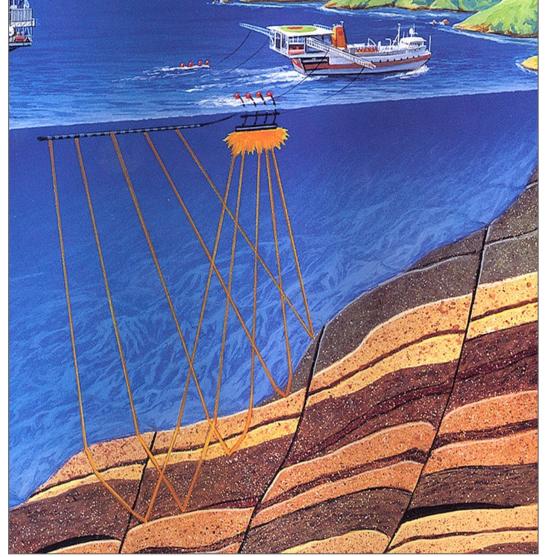




ACTIVE SEISMIC METHODS

Reflection seismic methods are more commonly used in geophysical exploration, as they require much shorter profiles and therefore less shot energy and have a much higher lateral resolution.

However, **reflection signals are much more complex** to detect and to analyse than refraction signals as they never arrive first, which implies time and labour intensive filtering and detection from a multitude of overlapping data. Moreover, the specific setup for reflection measurements requires more logistic preparation and personnel, which makes it generally a lot more expensive than refraction methods. It is the method of choice in hydrocarbon **exploration, as it can resolve structural details of a reservoir.** Supported by INTELLIGENT ENERGY EUROPE



GE Ge E C 7

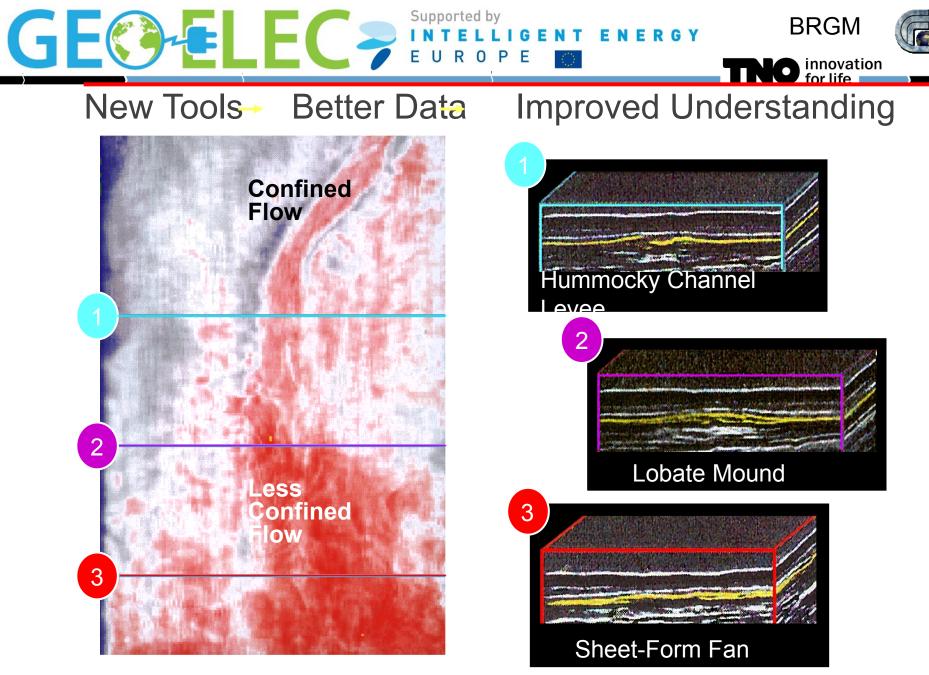
Seismic Imaging 3D Marine Data Acquisition

Silicon Graphics



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3D Seismic Image - Submarine Fan

Armentrout et al., 1996

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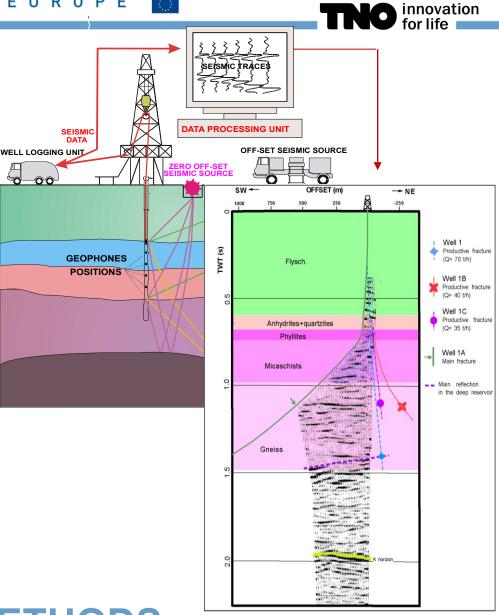
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Seismic signals generated and detected at the service are commonly restricted to horizontal or gently dipping reflectors. To detect and image vertical structures, vertical seismic profiling (VSP) was developed, which takes advantage of an existing well. VSP not only allow resolution of vertical reflectors such as faults but also provides highly reliable calibration tool for surface seismic and is useful in projects involving seismic anisotropy.

ACTIVE SEISMIC METHODS







Despite their clear advantages, especially resolution with depth, active seismic methods are not very common in geothermal exploration. One of the reasons why there are not widely used is that their cost often makes them difficult to fund for tight-budgeted geothermal projects especially where the geological complexity requires 3D arrays.

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In volcanic environment they are seldom used due to the too high noise and strong attenuation.

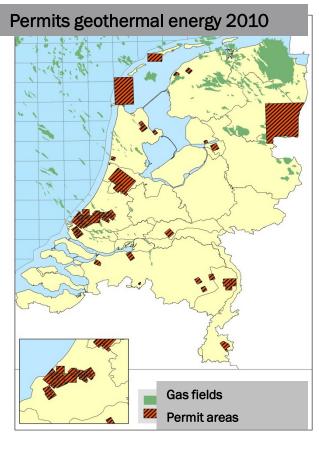


Reasons to develop ThermoGIS

Interest is booming

Currently over 100 permits granted

- Geological properties and uncertainties
- Independent analysis and information
- Overview potential areas and 'hot spots'
- > Performance assessment
- > Quickscan
- Realization of market opportunities

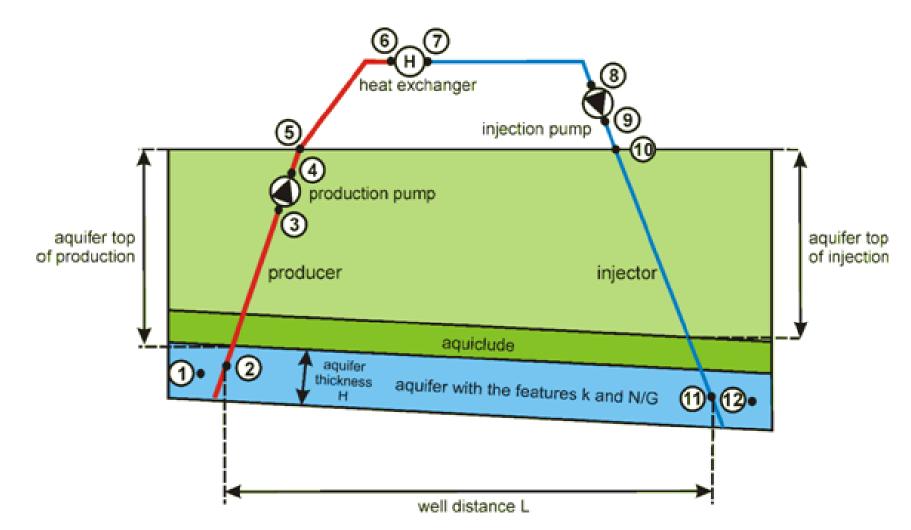


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Pre-feasibility analysis for heat production Schematic Doublet



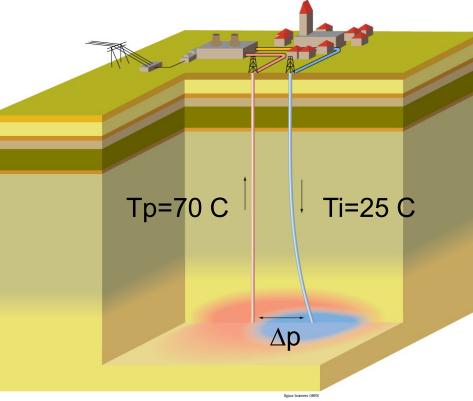


Doublet performance

 $E [MWth] = Q^* \Delta T * C_P$

Flow-rate Q

Permeability X thickness $Q = \Delta p \frac{2\pi k H}{\mu \left(\ln \left(\frac{L}{r_w} \right) - S \right)}$ Viscosity distance



∆p generated by pumps Which consume electricity

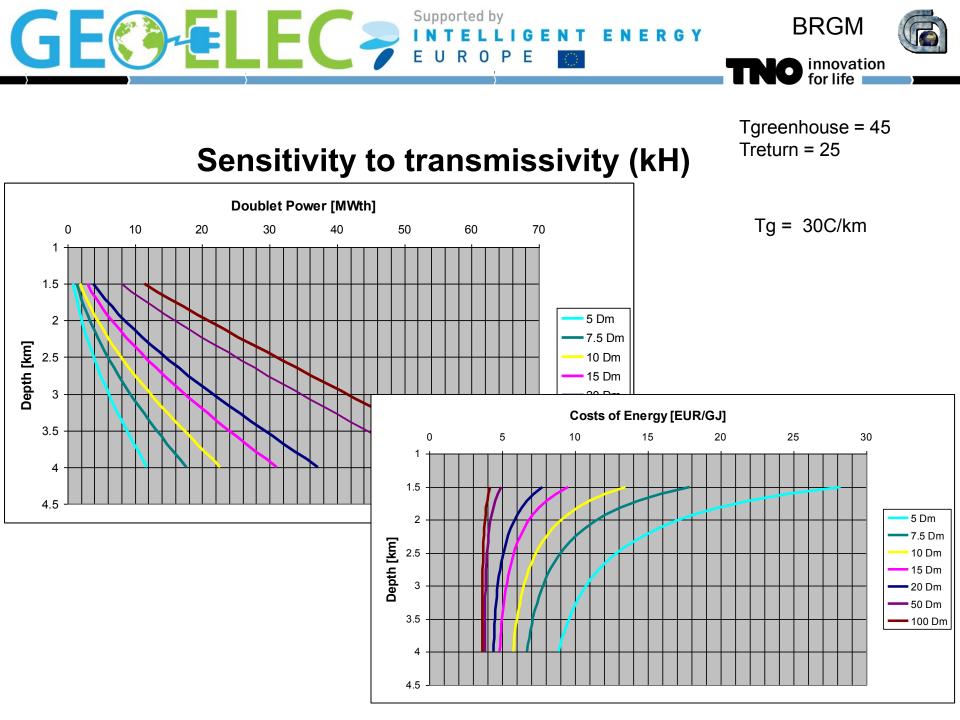
∆p is restricted by safety measures

 Δp at surface does not linearly lead to Higher flow rates (friction in tubes)

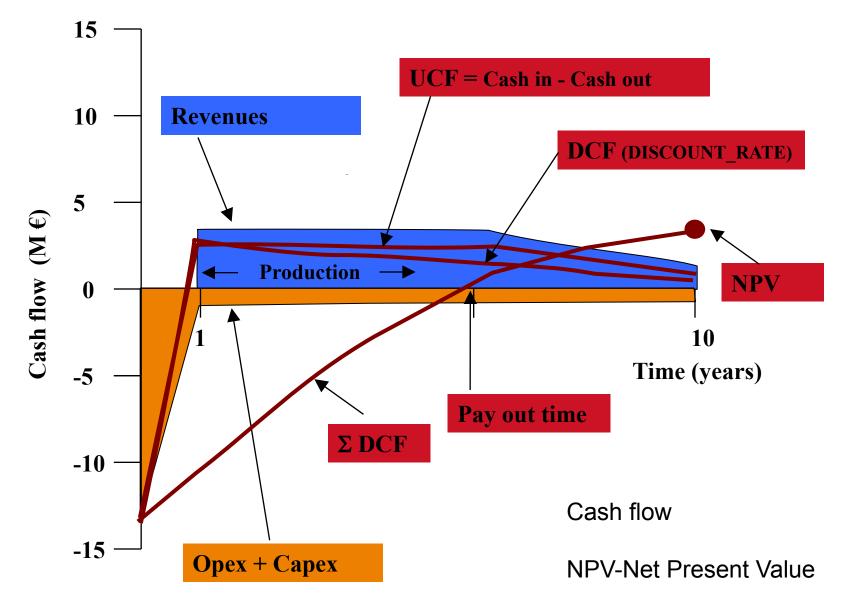


Potential estimates – for specific application areas

Geothermal direct heat Application	Min. Production Temperature	Min. production depth@ 30 °C/km	Re- injection temperatur e	Load factor	Heat demand/yr
Greenhouse	45°C	1200m	25°C	60%	12x10 ³ GJ/ha
Spatial	65°C	1900m	40°C	60%	25GJ/house









Levelized Cost of Energy

- > Discounted energy produced [MWh, GJ]
- Discounted cash out [EUR]
- LCOE = discouted cash-out / discounted energy produced



ThermoGIS ⁻ project



3D mapping reservoirs (aquifers)

- Depth, thickness and temperature → <u>Temperature</u>
- (Thickness) porosity, permeability → <u>Transmissivity</u>
- Uncertainties
- Potential energy

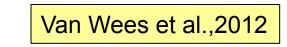
Pluymaekers et al.,2012

Bonte et al.,2012

Development ThermoGIS application

- Visualisation map
- Performance assessment tool
- Economic assessment tool

Kramers et al.,2012



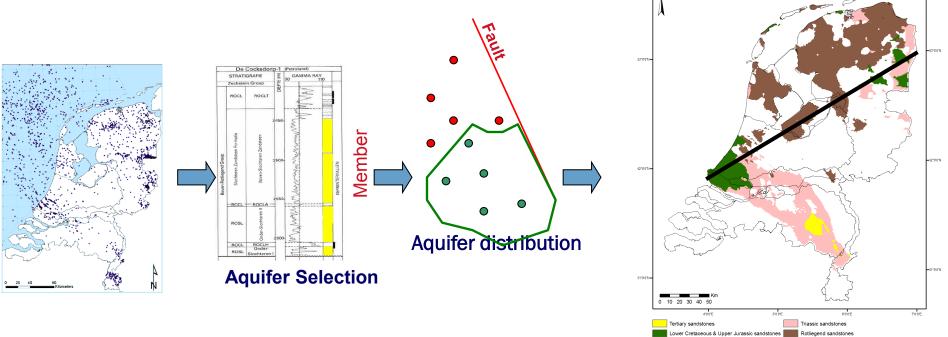


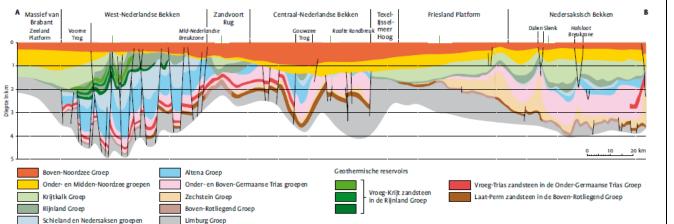
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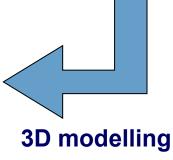


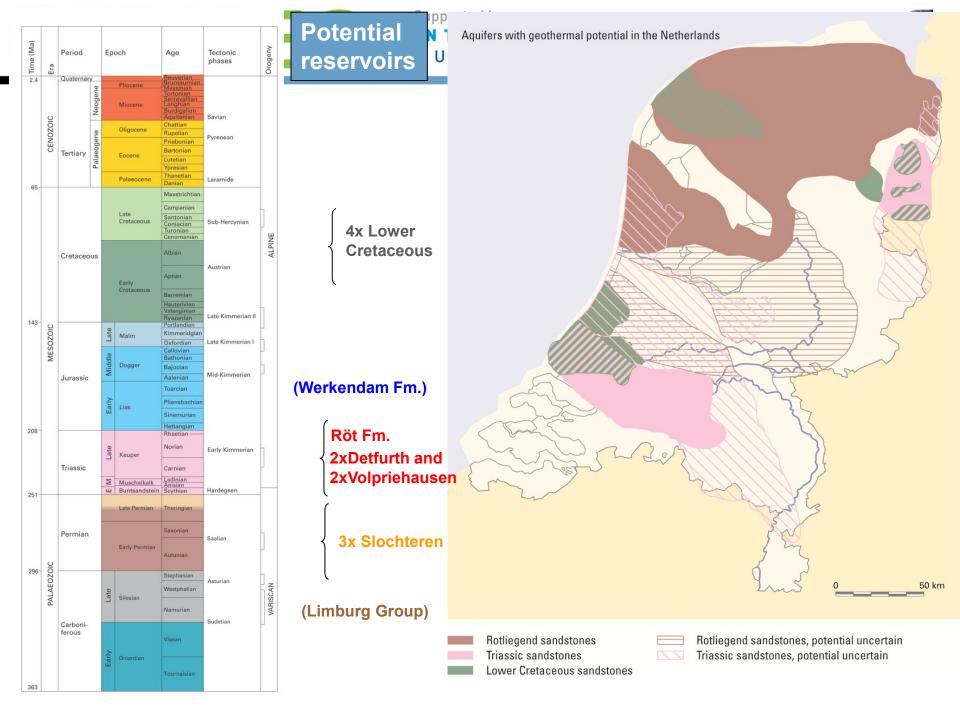
Aquifer depth and thickness mapping

GE()-€









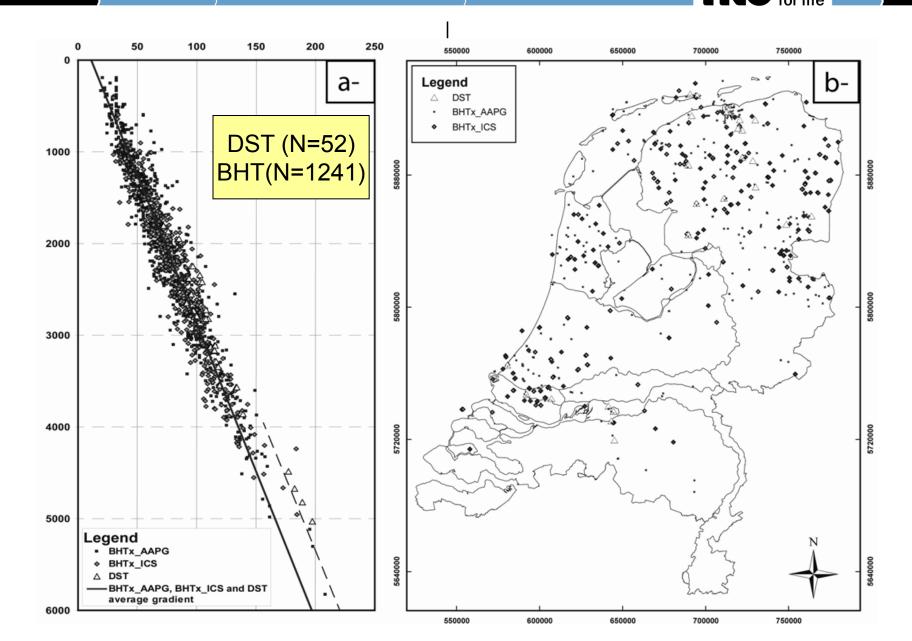
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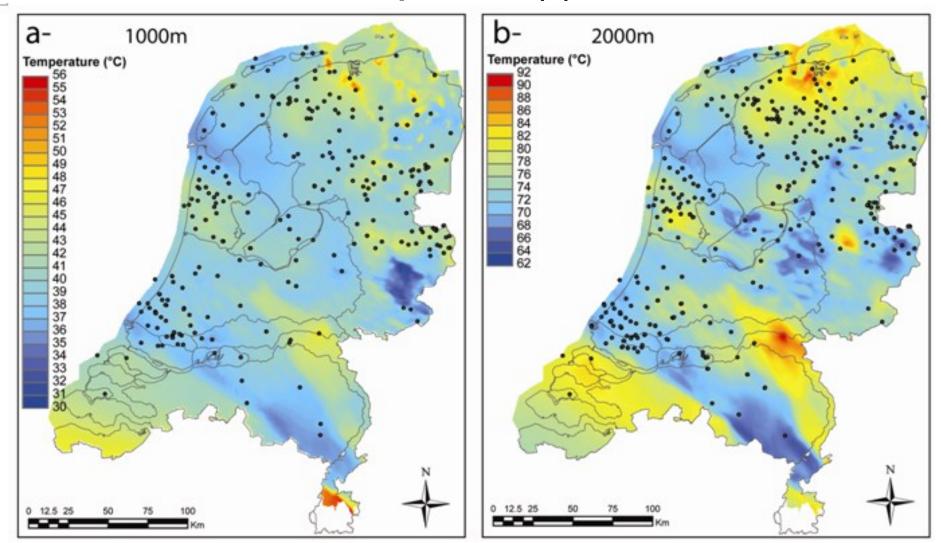
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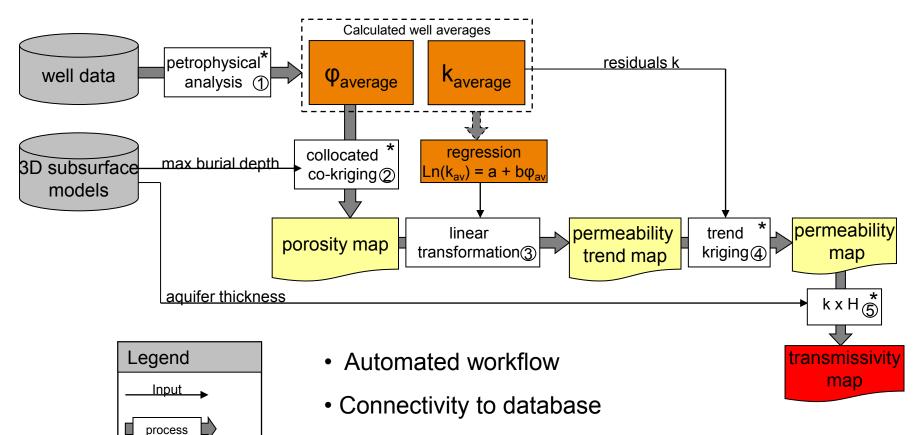
Results - temperature (1)





* marked by uncertainty

Property mapping



UROP

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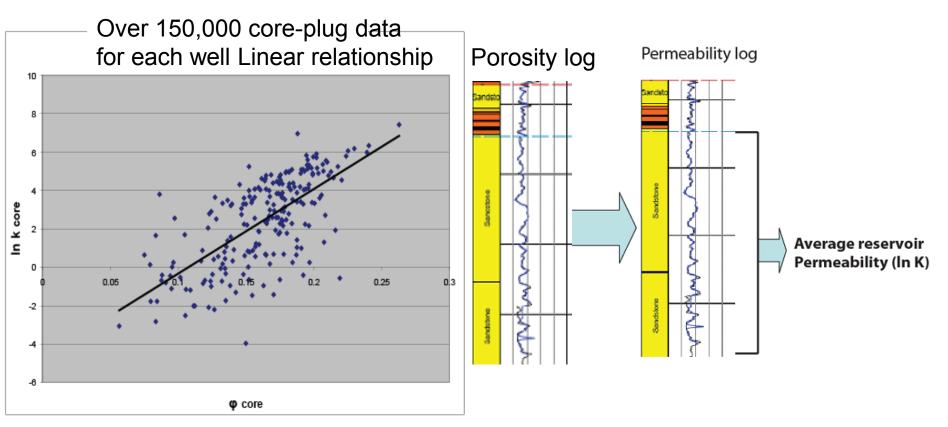
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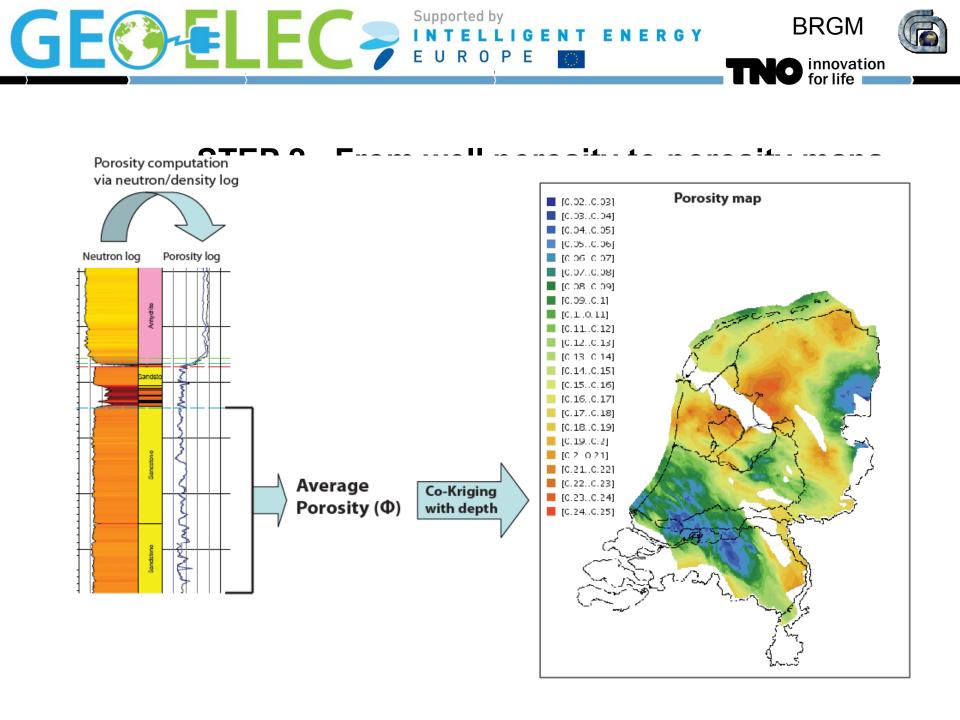
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Geostatisics for uncertainties



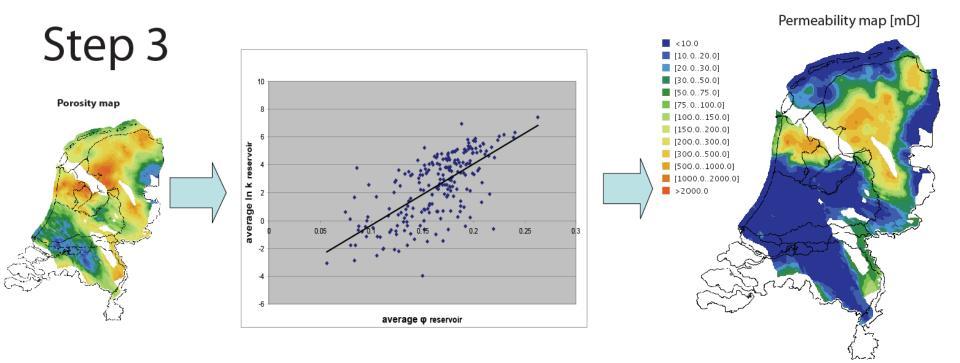
STEP 1: Porositv \rightarrow Permeability in wells

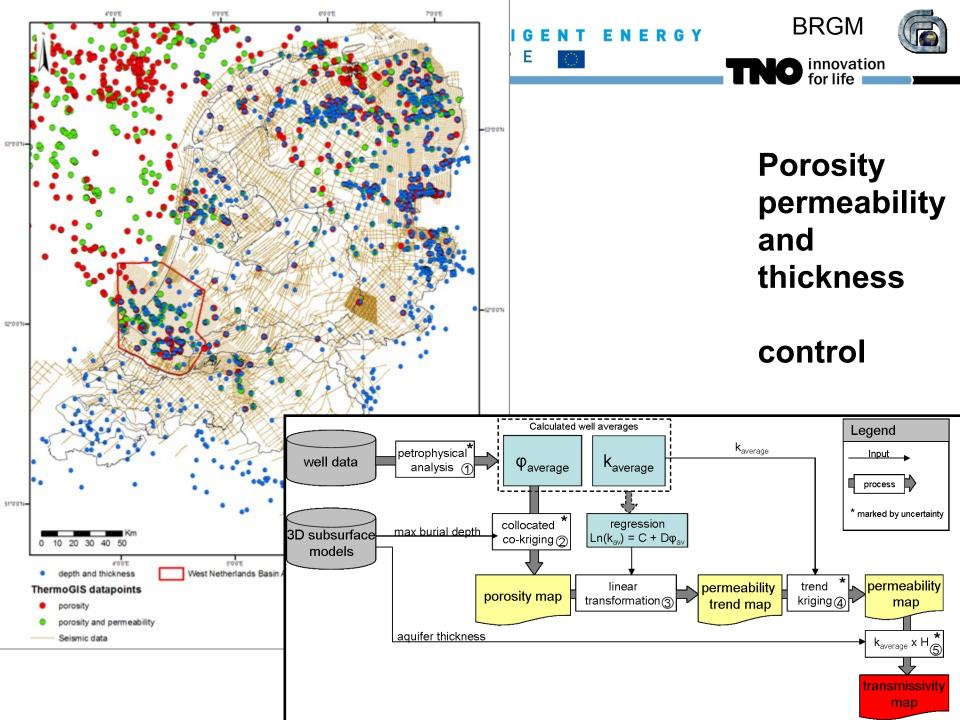






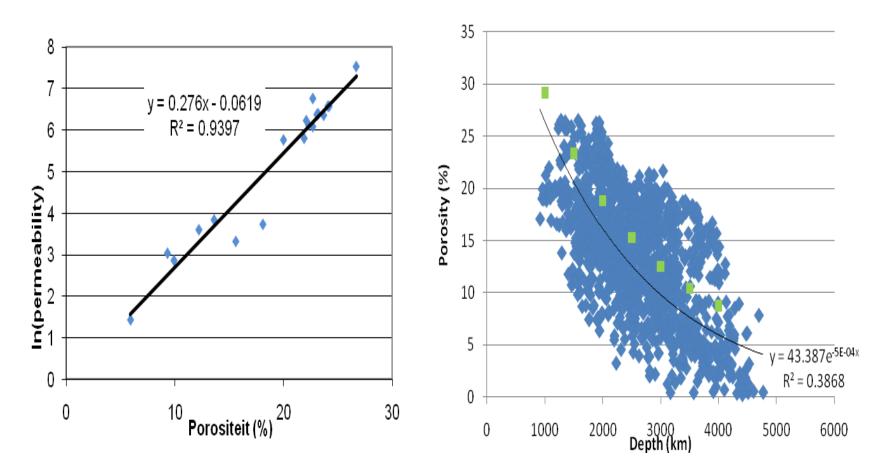
Average porosity – average permeability





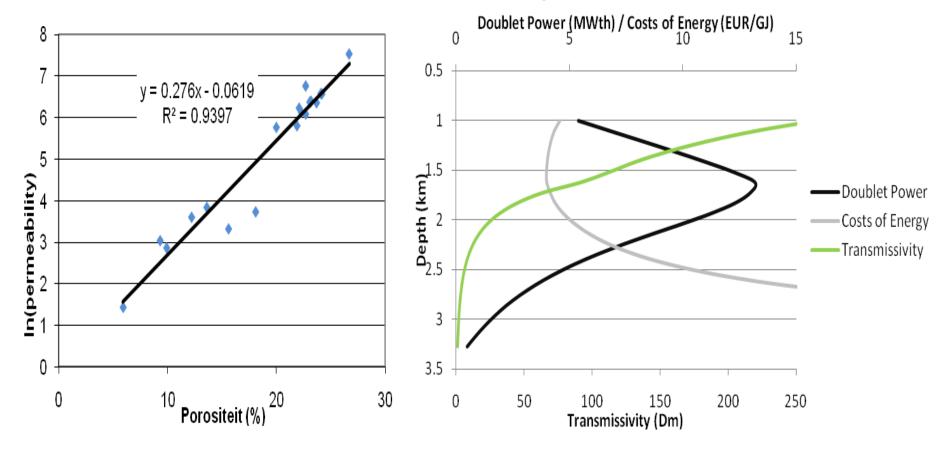


Permeability is determined using poro-perm relationship Porosity generally decreases with depth....



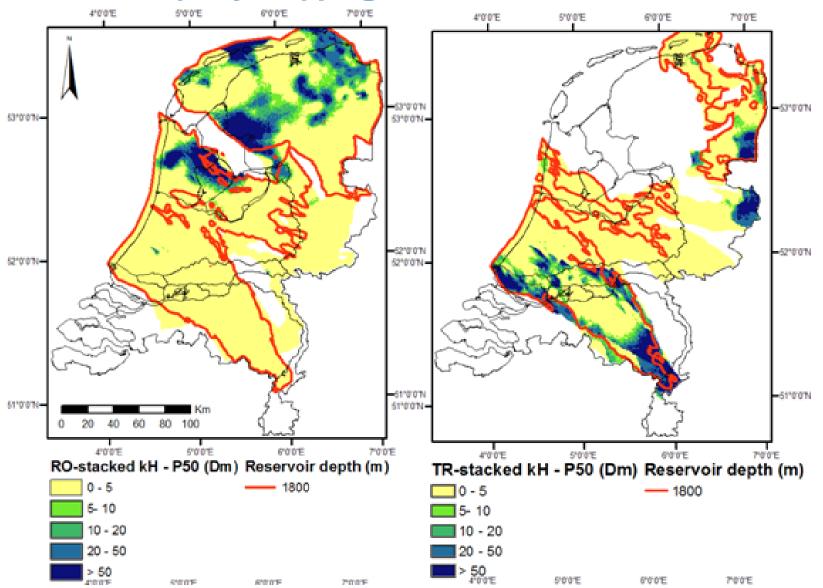


Performance is predicted to decrease with depth as a function of porosity



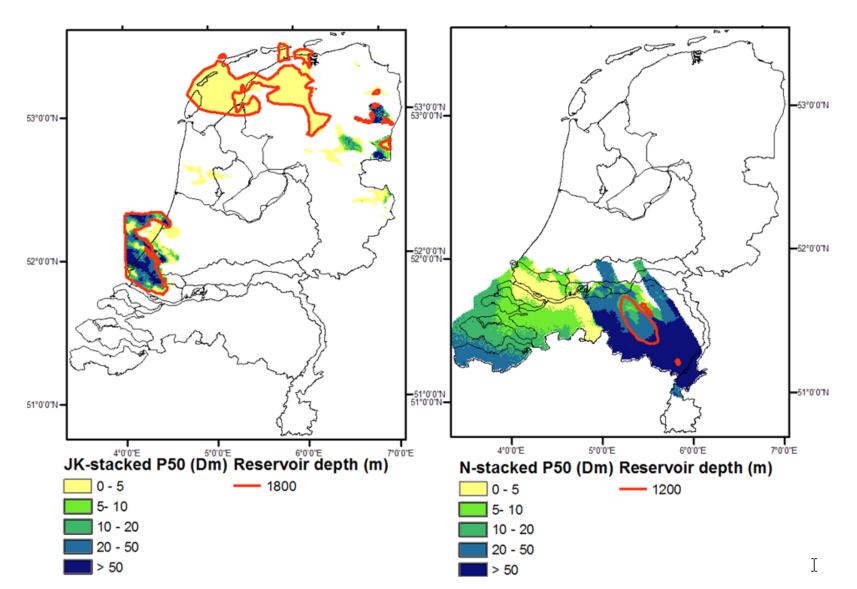


Property mapping and uncertainties



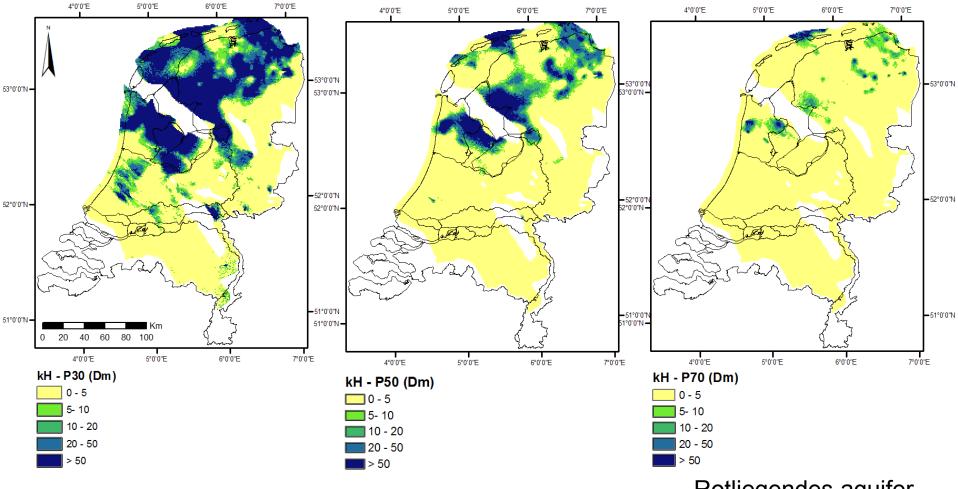


Property mapping and uncertainties (2)





Property mapping and uncertainties (3)



Rotliegendes aquifer

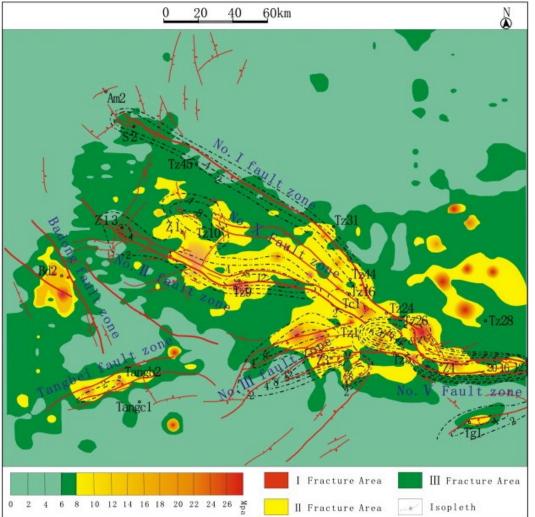


permeability in other sediments

- > Carbonates \rightarrow fracture related
- > Fracture assessment
 - > Well imaging of fractures
 - > Well flow tests
- ➤ Fracture related to tectonic deformation → mechanical models for natural fracture density variation related to position in structure



Carbonates permeability – learn from oil and gas

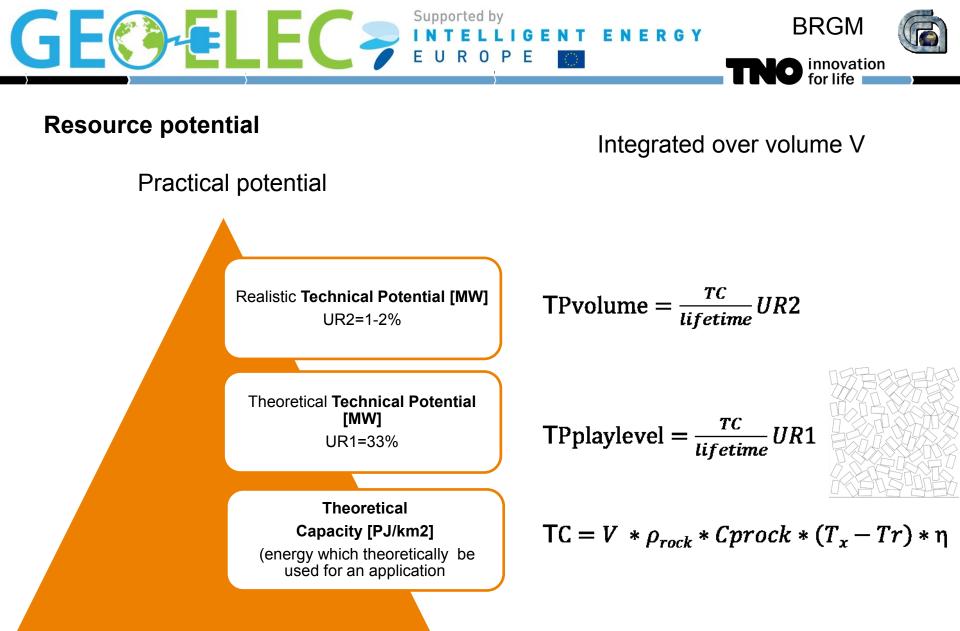


Ding et al., 2012

Journal of Petroleum Science and Engineering Volumes 86–87, May 2012, Pages 62–70

Ordovician carbonate reservoir fracture characteristics and fracture distribution forecasting in the Tazhong area of Tarim Basin, Northwest China

Predicted tensile stresses related To caledonian deformation



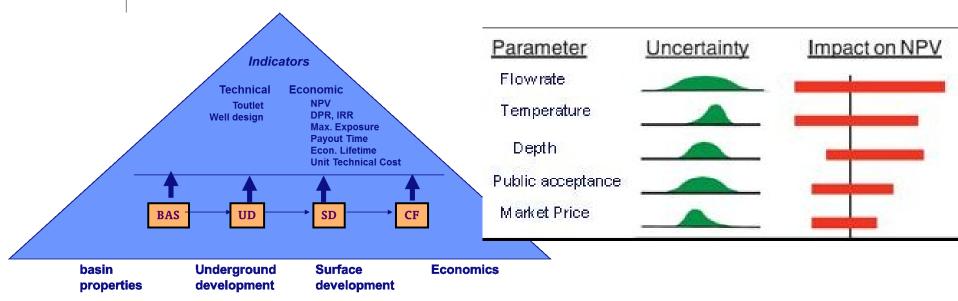
Theoretical potential

Beardsmore et al., 2011. philosophy used In IPCC and IEA roadmap

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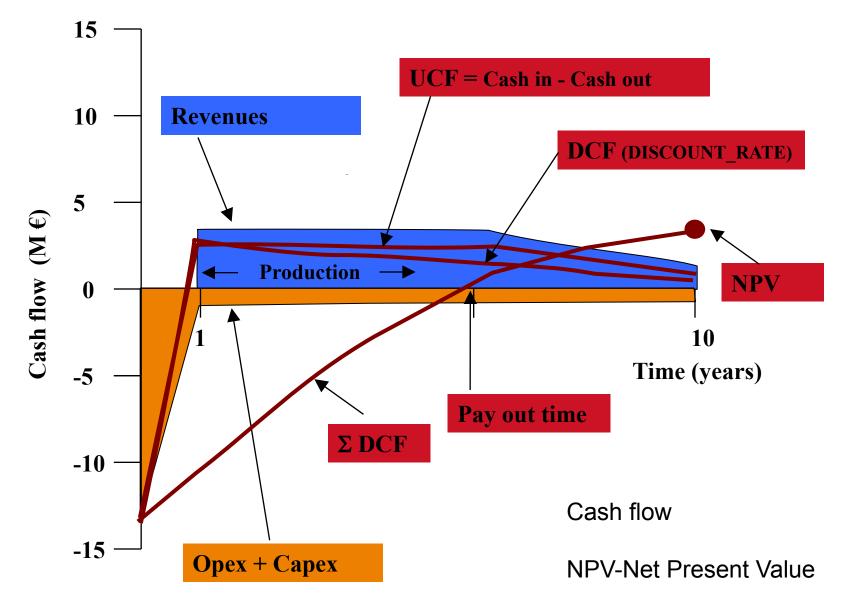
Resource Potential: how to achieve transparant framework

- Evaluate transparant Key Performance Indicators
 - Net Present Value, Levelized Cost of Energy
 - Evaluate with fastmodels for techno-economic performance
 - Use MC sampling evaluate risk and upside in reward



driving philosophy is to trade-off accuracy for completeness

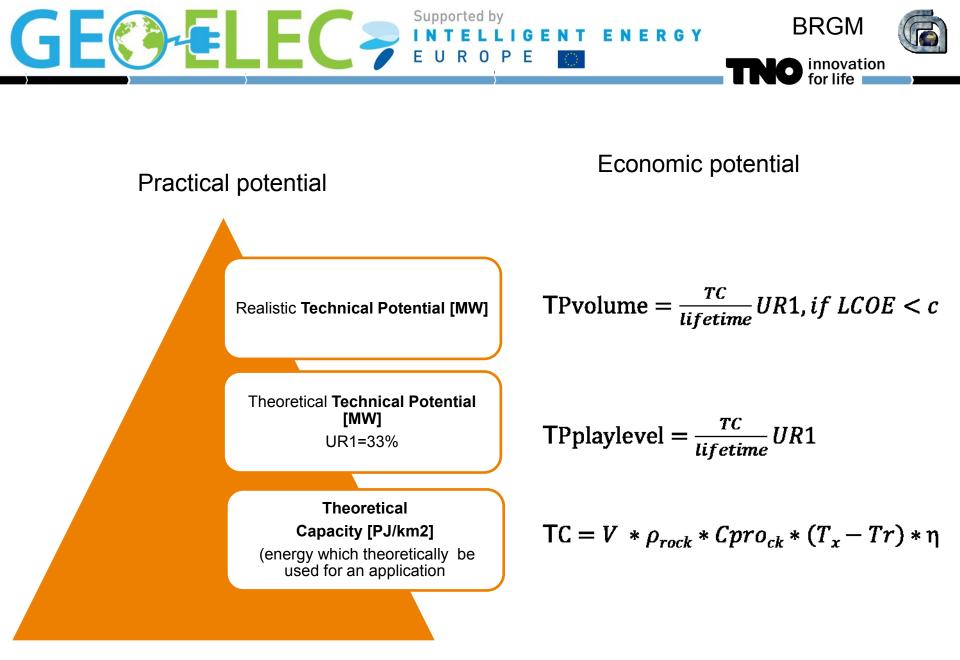






Levelized Cost of Energy

- > Discounted energy produced [MWh, GJ]
- Discounted cash out [EUR]
- LCOE = discouted cash-out / discounted energy produced

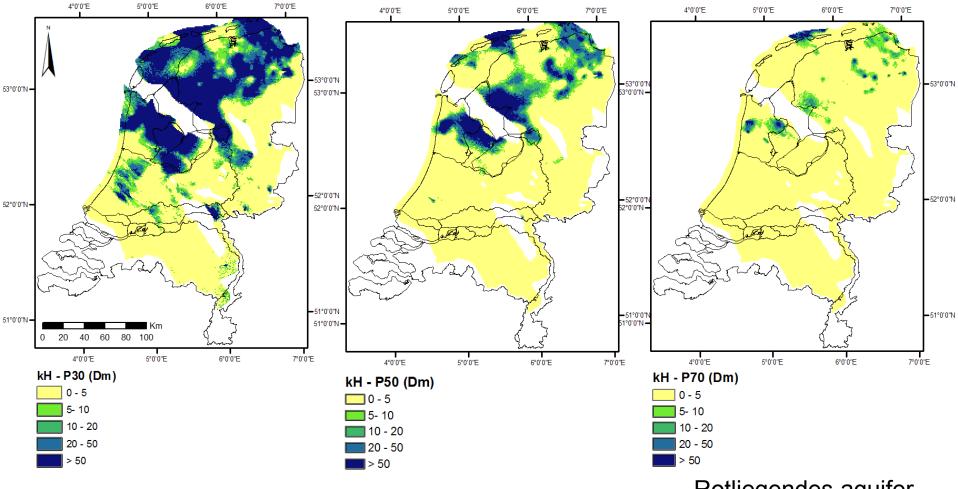


Theoretical potential

Used here



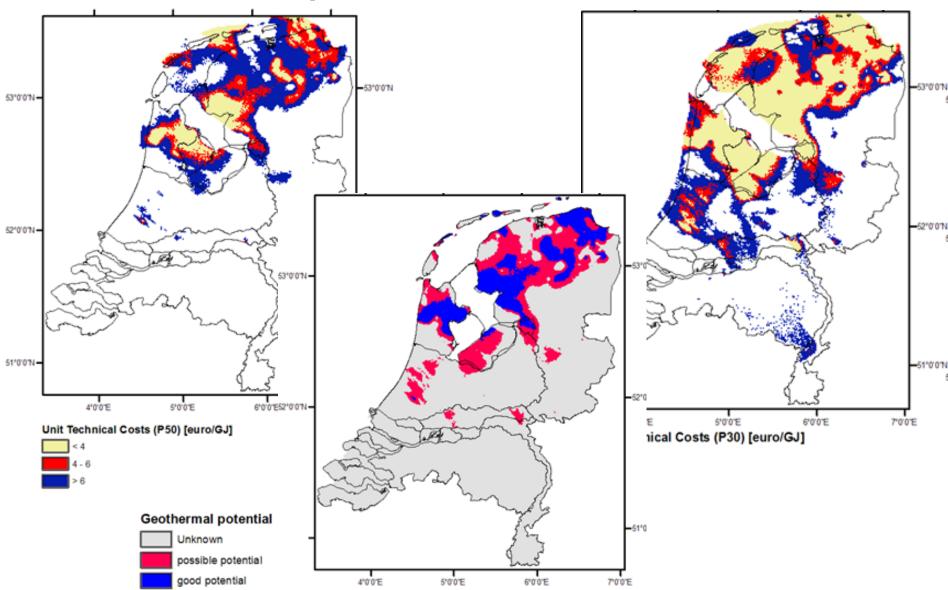
Property mapping and uncertainties (3)



Rotliegendes aquifer

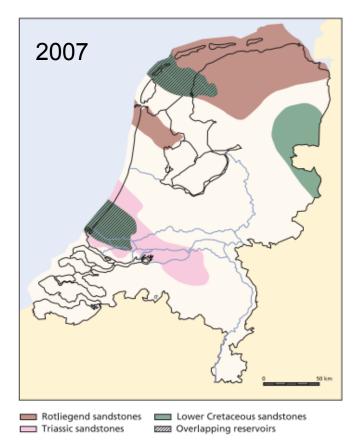


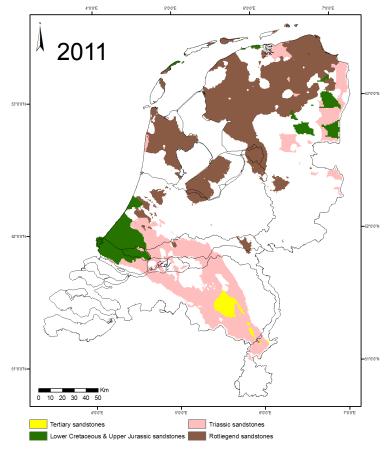
Potential Map





potential map vs earlier assessments





Mapped stratigraphic areas

White: insufficient amount of data or other stratigraphic units (like Carboniferous)

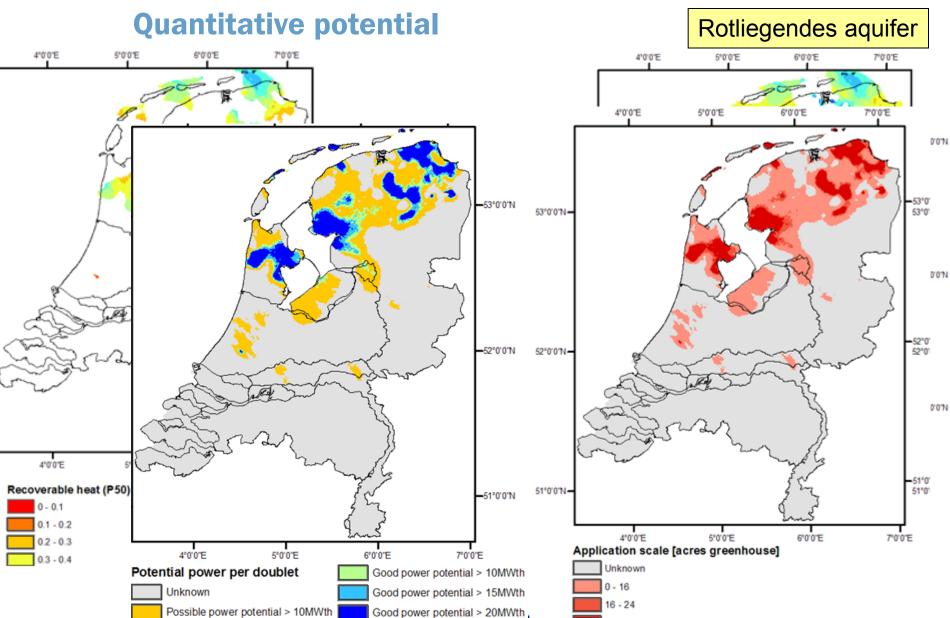


24 - 32

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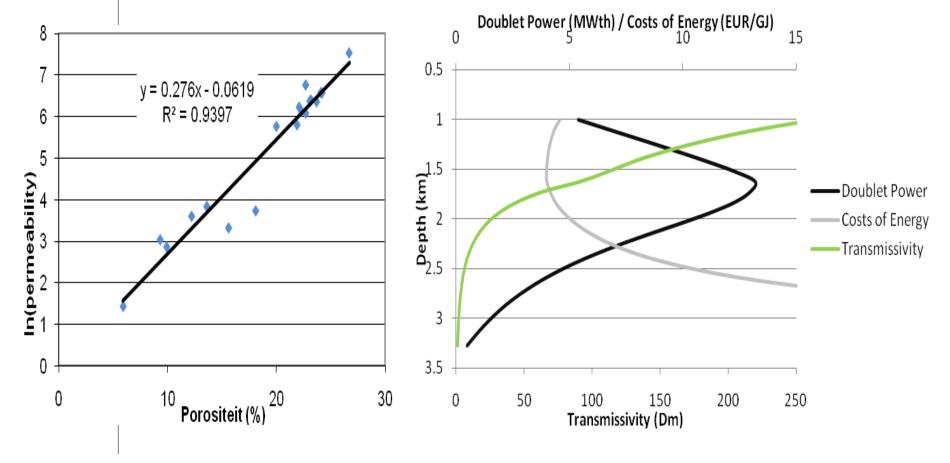


INTEL

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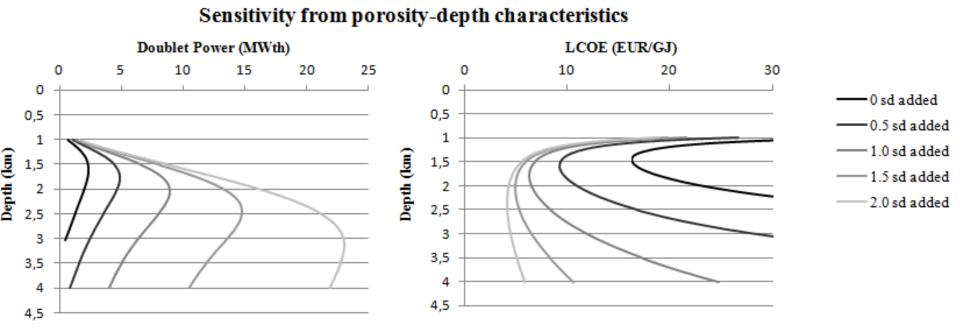
Performance is predicted to decrease with depth as a function of porosity

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Analysis of sensitivity to typical dutch permeability trends





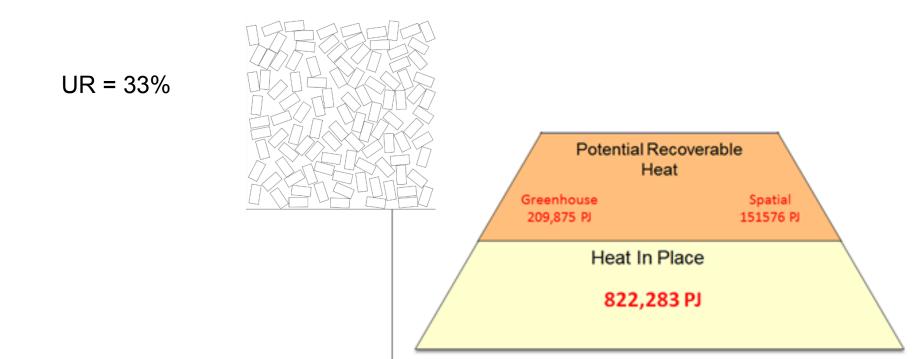
Heat In Place [PJ km-2] Starting point at the base of the pyramid is Heat In Place (HIP) in PJ/km2. This is the heat content of the reservoir (cf. Muffler and Cataldi, 1978). HIP is the maximum theoretically extractable heat in the aquifer.

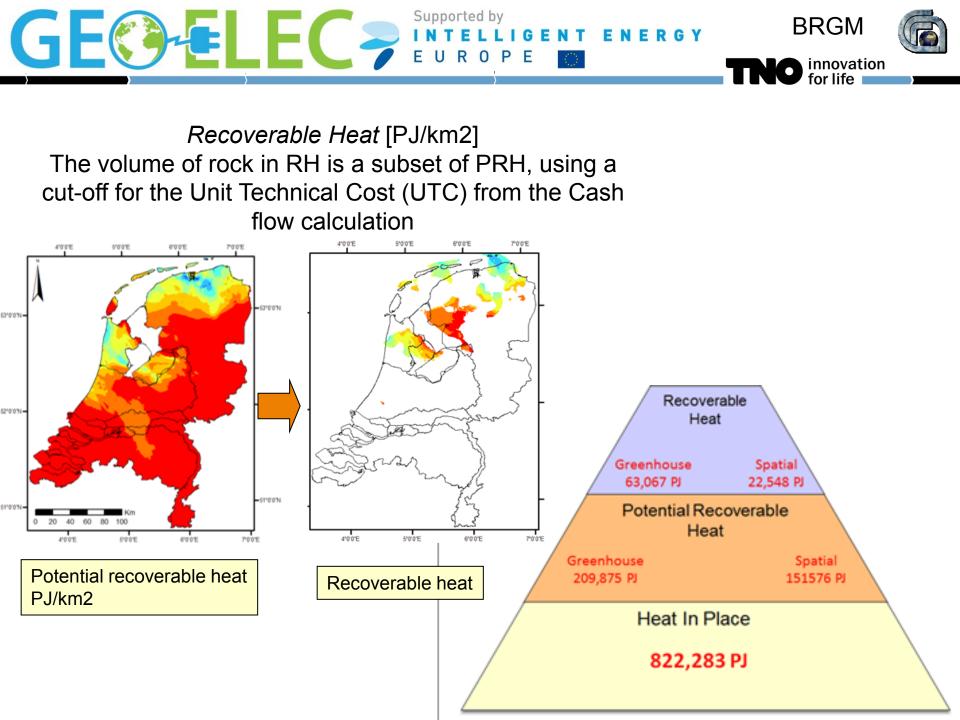
Heat In Place

822,283 PJ



Potential Recoverable Heat [PJ/km2] or Technical potential [PRH/30 yr] The next level of the pyramid is the Potential Recoverable Heat (PRH). This is the heat which can be recovered from the reservoir, unconstrained by economic limitations and irrespective of flow properties.







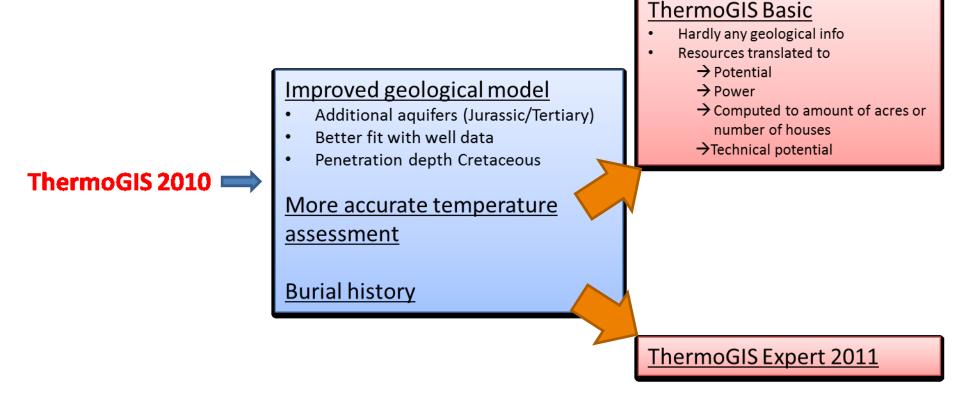


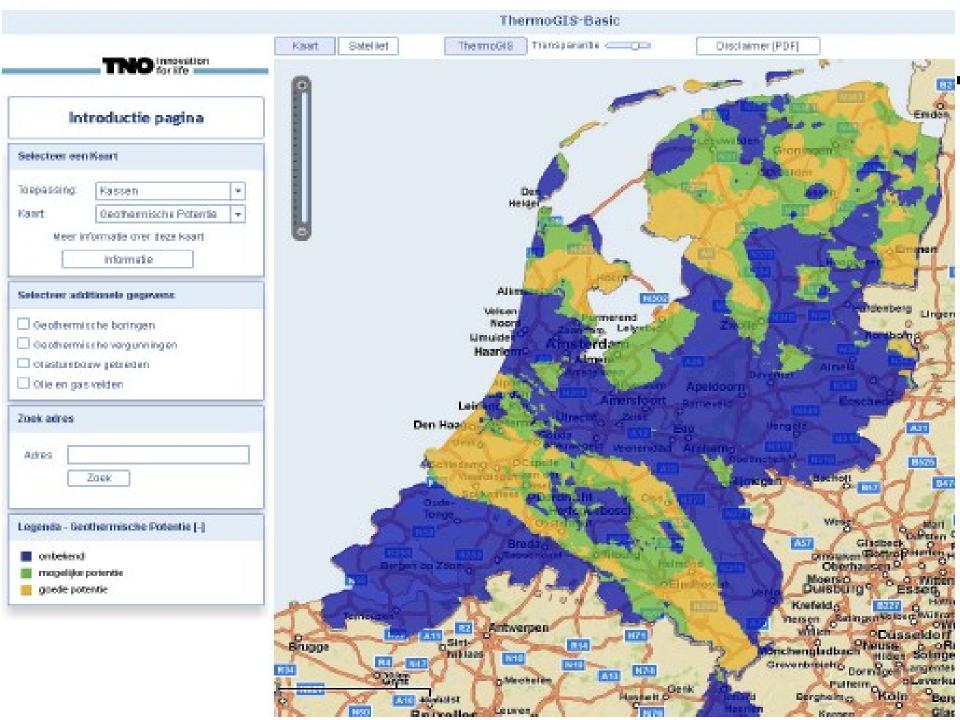
for life

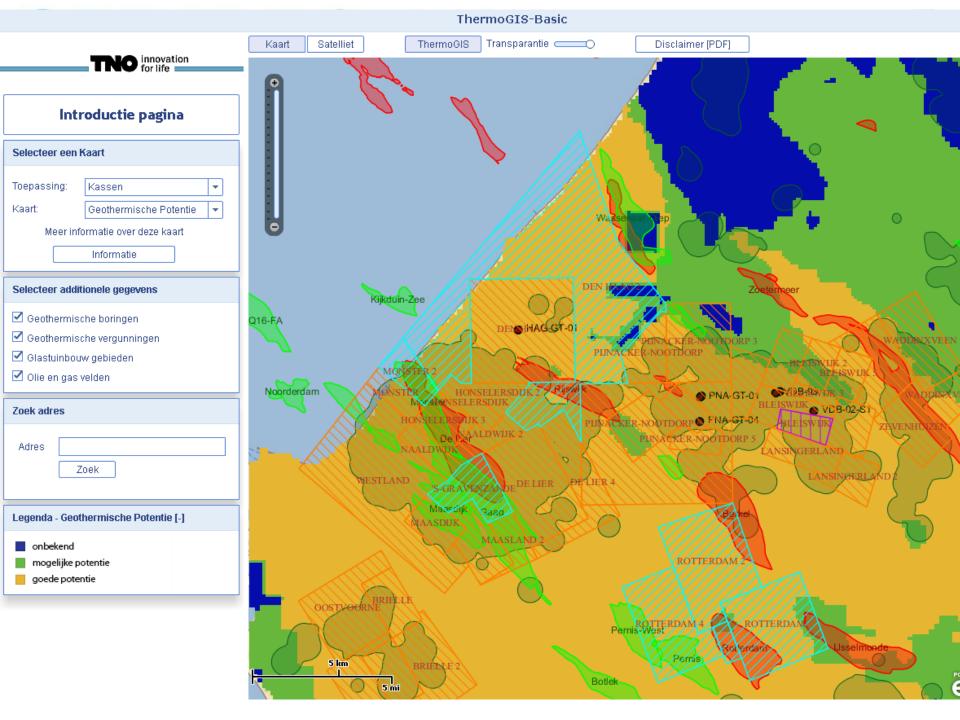


Thermc |S evolution

GF





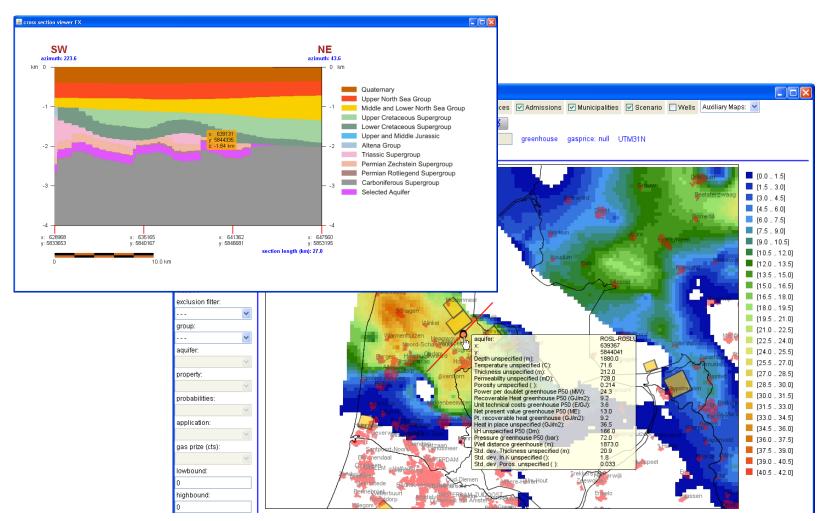


ThermoGIS Expert: Geological properties

GE

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THERMOGIS[™] doublet thermal Power [MWth], site specific information



Result screen "Geothermal Power Program"

Probability Density Function of Geothermal Power

Probability plot: retrieved geothermal energy (MW)

